

Investigating the Impact of Snowpack Photodenitrification on Polar Atmospheric Chemistry Using Results from a Snowpack Radiative Transfer Model



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Maria Zatkan
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Snowpack Photochemistry

GOAL: Incorporate process-based study of snowpack photodenitrification into a global chemical transport model (GEOS-Chem) to determine spatial redistribution of nitrate

-collaborators using GEOS-Chem adjoint model to determine sources of nitrate to Antarctica

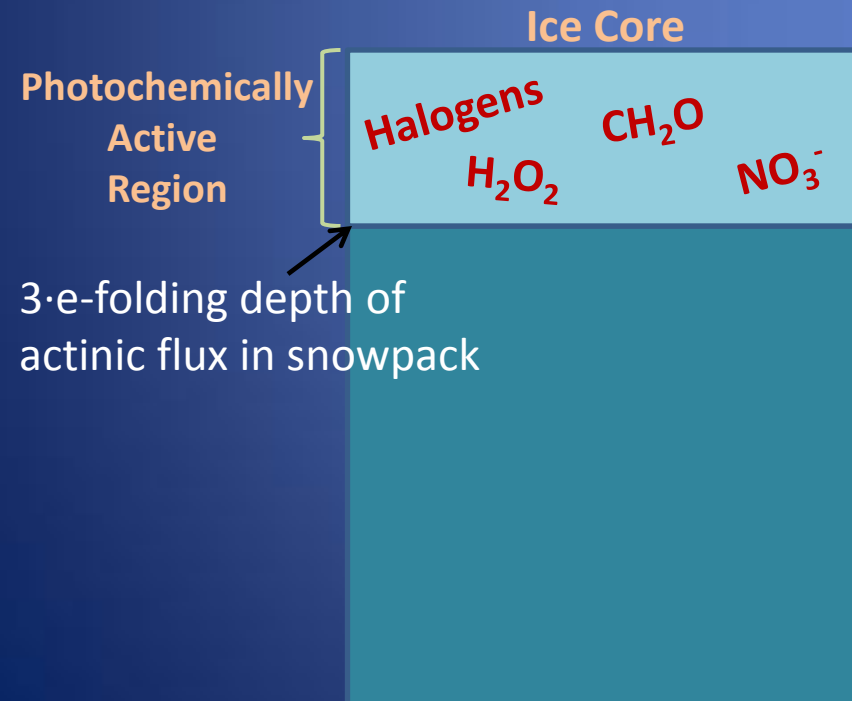
This Study

-Snowpack radiative transfer model with updated optical properties of ice in UV used

-Sensitivity of snowpack actinic flux profile to snowpack physical and optical properties explored

- τ_{NO_x} in snowpack interstitial air against physical (τ_{escape}) and chemical (τ_{chemical}) loss determined

-parameterization for actinic flux in snowpack developed for use in large scale models



Snowpack Radiative Transfer Model

Wolff et al. [2002]

Snowpack Radiative Transfer Model
[Grenfell, 1991]



e-folding depth of actinic flux in snowpack of 3.7 cm at Neumayer

This Study

Snowpack Radiative Transfer Model
[Grenfell, 1991]

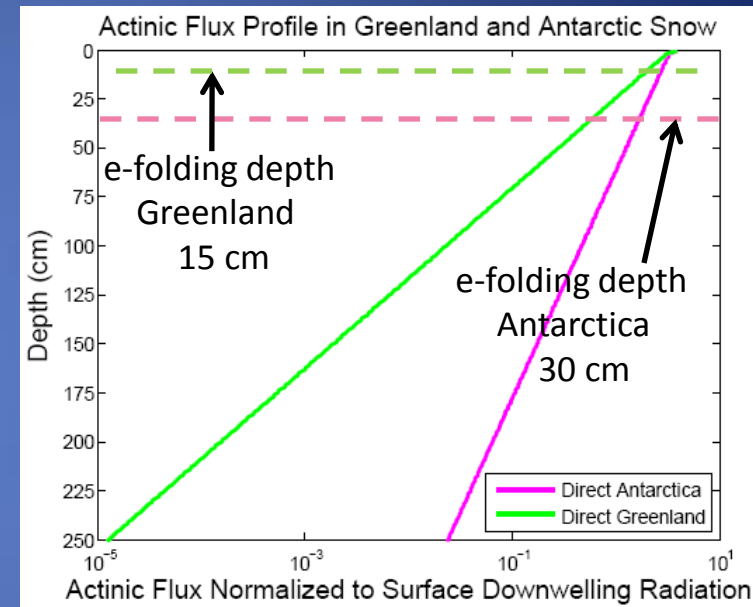


-Updated optical properties of ice [Warren and Brandt, 2008]
-Inclusion of soot and dust as dominant absorbers



e-folding depth of actinic flux in snowpack of 30 cm at Neumayer

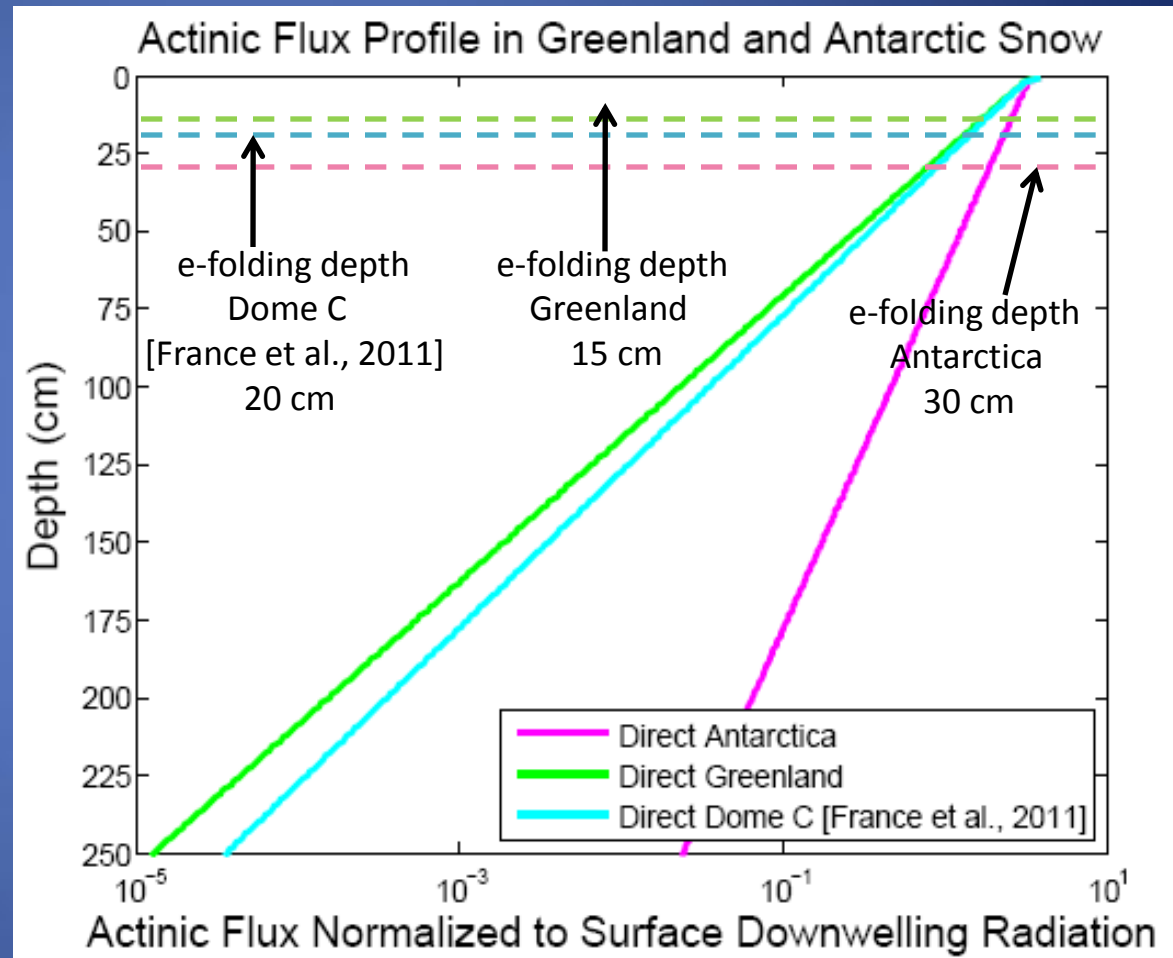
Profile of Actinic Flux in Snowpack



Snowpack Radiative Transfer Model

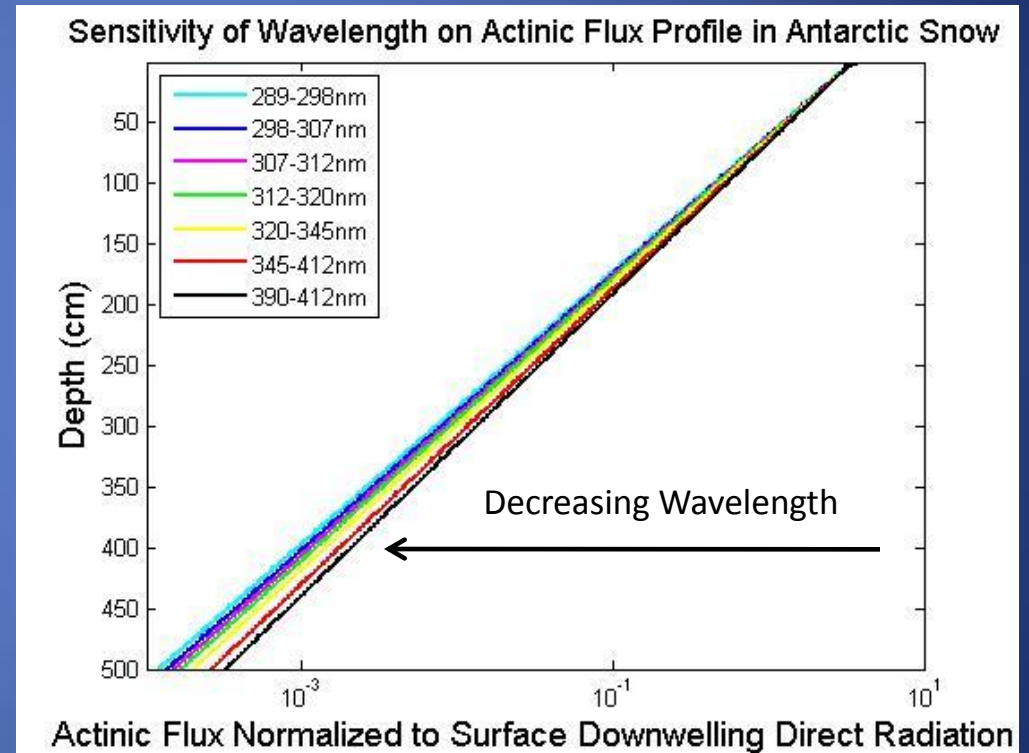
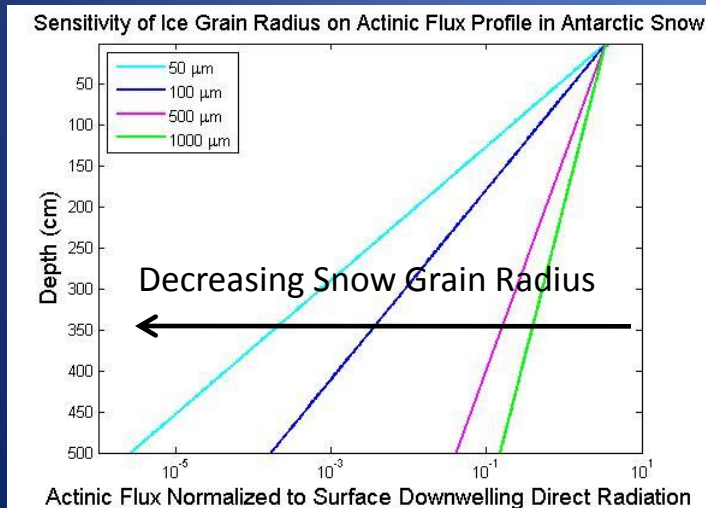
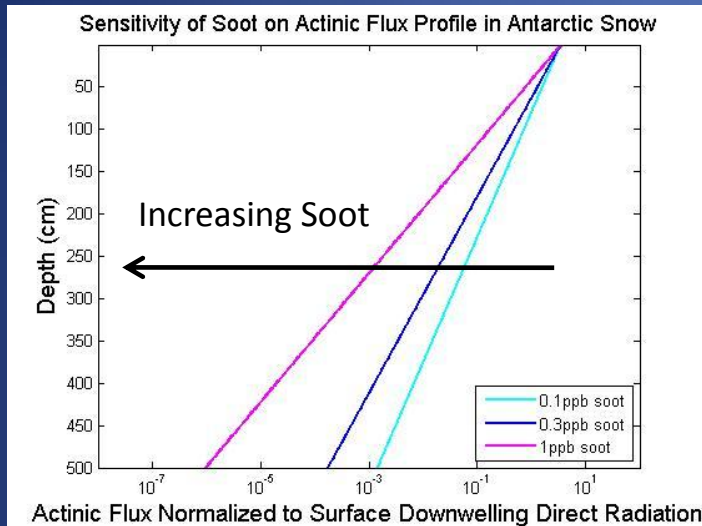
Our parameterizations agree with recent e-folding depth observations from France et al. [2011] when same snow conditions used at Dome C

Soot concentrations observed by France et al. [2011] (~ 3 ppb) greater than concentrations in clear air sectors in Antarctica (0.3 ppb) [Warren and Clarke, 1990]



Actinic Flux Profiles in Snowpack

Diffuse vs. direct radiation, solar zenith angle, wavelength, effective snow grain radius [Hansen and Travis, 1974], density, soot concentration and dust concentration



Actinic flux in deep snowpack most influenced by soot concentration and effective snow grain radius

Will NO_x produced at a given depth in snowpack be released into atmosphere?

Determine ventilation depth by comparing the lifetime (τ) of NO_x against:

Diffusion

Wind pumping

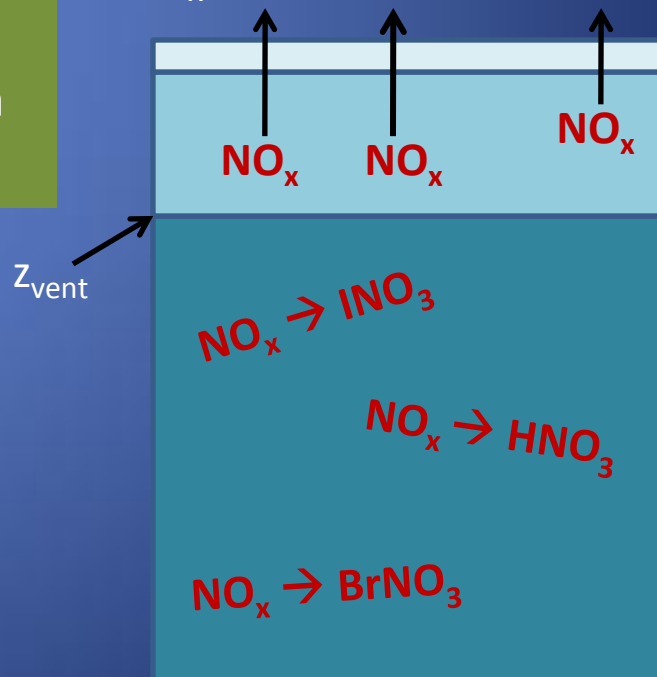
Conversion to HNO_3

τ_{escape}

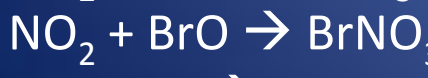
τ_{chemical}

Ventilation depth (z_{vent}):
depth below which NO_x produced is unlikely to escape to atmosphere before reacting to form HNO_3 .

NO_x Escape vs. Conversion



Conversion of NO_x into HNO_3 , BrNO_3 , and INO_3 in snowpack prevents NO_x ventilation to atmosphere.



hydrolysis



NO_x Fluxes

Burley and Johnston [1992]

$$J_{\text{NO}_3^-} = \iint \sigma(\lambda) \cdot \phi(T, \text{pH}) \cdot I_0(\lambda, z) d\lambda dz$$

Chu and Anastasio [2003]

Our parameterizations based upon Grenfell [1991] and Wild et al. [2000]

$$F_{\text{NO}_x} = J_{\text{NO}_3^-} [\text{NO}_3^-]$$

Dibb et al. [2004], [2007]
 Mulvaney et al. [1998]
 Honrath et al. [2002] x 6%

Thomas et al. [2011]

NO _x Flux (molecules cm ⁻² s ⁻¹)			
Locations	This Study	Observations	Reference
Neumayer (coastal)	7.1x10 ⁷ -2.9x10 ⁸	1.3-2.5x10 ⁸	Jones et al., 2000 Jones et al., 2001
South Pole (continental)	1.0-1.8x10 ⁸	2.2-3.8x10 ⁸	Davis et al., 2004 Oncley et al., 2004
Halley (coastal)	9.2x10 ⁷ -3.7x10 ⁸	1.7-3.8x10 ⁸	Jones et al., 2007 Bauguitte et al., 2009
Summit (continental)	9.5x10 ⁷ -1.8x10 ⁸	2.5x10 ⁸	Honrath et al., 2002

NO_x fluxes calculated in this study are compared to NO_x flux observations at South Pole, Neumayer, Halley, and Summit

Snowpack Actinic Flux Parameterizations

For direct radiation at the surface:

$$\left[\frac{I_o(\lambda, z=0)}{F_{inc}(\lambda_{\downarrow})} \right]_{direct} = \left[\frac{0.577 + \mu_o}{0.577 \cdot \mu_o} \right] \cdot Corr(\mu_o)$$

correction factor
because no exponential
decay in top 2 cm

solar zenith angle

2 cm

For direct radiation below z_{ref} :

$$\left[\frac{I_o(\lambda, z \geq z_{ref})}{F_{inc}(\lambda_{\downarrow})} \right]_{direct} = G(z_{ref}, \mu_o) \cdot e^{-0.60 \cdot c\overline{\omega}_{eff}^{\frac{1}{2}} \cdot Kext_{tot} \cdot (z - z_{ref})}$$

$$G(z_{ref}, \mu_o) = 3(0.577 + \mu_o) \cdot e^{-0.60 \cdot c\overline{\omega}_{eff}^{\frac{1}{2}} \cdot Kext_{tot} \cdot z_{ref}} \cdot Corr(\mu_o)$$

For diffuse radiation:

$$\left[\frac{I_o(\lambda, z)}{F_{inc}(\lambda_{\downarrow})} \right]_{diffuse} = 3.831 \cdot e^{-0.60 \cdot c\overline{\omega}_{eff}^{\frac{1}{2}} \cdot Kext_{tot} \cdot z}$$

Actinic Flux

$$I_{o_{combined}}(\lambda, z) = \left\{ \left[\frac{I_o(\lambda, z)}{F_{inc}(\lambda_{\downarrow})} \right]_{diffuse} \cdot (f_{dif}) + \left[\frac{I_o(\lambda, z)}{F_{inc}(\lambda_{\downarrow})} \right]_{direct} \cdot (1 - f_{dif}) \right\} \cdot [F_{inc}(\lambda_{\downarrow})]_{tot}$$

Extinction coefficients and
co- albedo for single scattering

$$Kext_{snow} = \frac{3 Q_{ext} \cdot \rho_{snow}}{4 r_e \cdot \rho_{ice}}$$

[soot] in snow

$$Kext_{soot} = \frac{\beta_{soot} \cdot L_{soot} \cdot \rho_{snow}}{c\overline{\omega}_{soot}}$$

[dust] in snow

$$Kext_{dust} = \frac{\beta_{dust} \cdot L_{dust} \cdot \rho_{dust}}{c\overline{\omega}_{dust}}$$

$$c\overline{\omega}_{eff} = \frac{c\overline{\omega}_{snow} \cdot Kext_{snow} + c\overline{\omega}_{soot} \cdot Kext_{soot} + c\overline{\omega}_{dust} \cdot Kext_{dust}}{Kext_{tot}}$$

$$Kext_{tot} = [Kext_{snow} + Kext_{soot} + Kext_{dust}]$$

The equation for actinic flux
in snowpack can be integrated
over wavelength and depth

Conclusions and Future Research

- The calculated e-folding depth of actinic flux is ~30 cm in Antarctic snowpack and ~15 cm in Greenland snowpack.
- Ventilation depths are most sensitive to sastrugi dimensions, [BrO], and [IO]. In absence of a ventilation depth, use 3-e-folding depth. (effect $z_{\text{vent}} = 90$ cm in Antarctica and 45 cm in Greenland)
- We calculate a wide range of NO_x fluxes based upon variations in ventilation depths and $[\text{NO}_3^-]$ in snow which are in agreement with observations in Antarctica and Greenland
- We have developed simple and broadly applicable equations to calculate depth dependent actinic flux in snowpack. These equations can be incorporated into global models and adjusted to represent all snowpack types by varying relevant parameters (e.g. soot, dust, snow density).
- Our next step is to incorporate our methods and results into the GEOS-Chem global chemical transport model to investigate the impacts on polar nitrogen and oxidant budgets.